

# Advanced Material Models for Seismic Simulations using ADER-DG

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ENERXICO



TUM Uhrenturm

# Motivation



Seismic wave propagation for applications in the energy sector

- Not just elasticity
- Cracked or damaged rock
- Interaction with a fluid phase

⇒ anisotropic and poroelastic materials



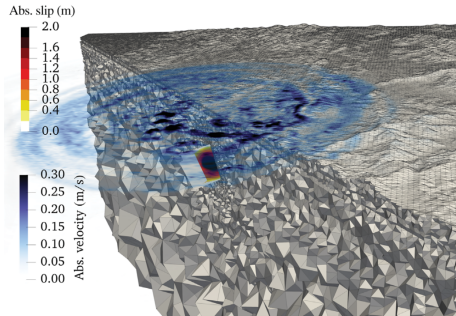
In particular we are interested in the interplay of induced earthquake rupture, wave propagation, complex topography and poroelastic effects.

# Earthquake simulations

Use SeisSol to solve the elastic wave equation:

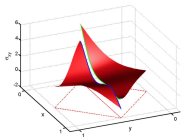
- $\partial_t q + A\partial_x q + B\partial_y q + C\partial_z q = 0$
- $q$  collects stresses and velocities

(Palgunadi et al. 2020):  
 "Dynamic Fault Interaction  
 during a Fluid-Injection-  
 Induced Earthquake: The  
 2017 Mw 5.5 Pohang Event"



# Discontinuous Galerkin method

- Discretize  $\Omega$  in tetrahedrons
- Expand the solution in terms of polynomials  $q_p(t, \vec{x}) = Q_{pI}(t)\phi_I(\vec{x})$ .
- Multiply the PDE with an element local test function  $\psi$  and integrate by parts
- Use numerical fluxes to exchange informations between elements.



# High-Order time stepping

Can be combined with time stepping like Runge-Kutta or Arbitrary DERivatives ansatz.

- Expand solution in time as a Taylor series around  $t^n$  to predict solution
- Use space derivatives at time  $t^n$  to get the time derivatives with the Cauchy-Kowalevski procedure
- Use fluxes to correct the solution

⇒ Achieve same convergence order in space *and* time

# HPC Optimizations

## Parallelization

- Element local discretization with DG
- Mesh partitioning based on workload estimate
- Exchange values at partion boundaries

## Node-level performance

- Update scheme is a sequence of tensor contractions
- Use YATeTo<sup>1</sup> to map the tensor operations to GEMMs  
( $C = \alpha AB + \beta C$ )
- Use architecture specific backends (like `libxsmm`) for optimized code

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<sup>1</sup>(Uphoff and Bader 2020)

## Model extensions – Anisotropy

Anisotropic material: directional dependent material behaviour

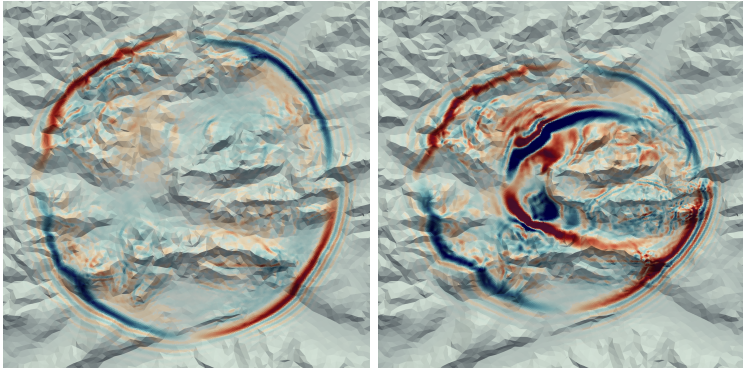
In seismology: layered or cracked media

For SeisSol: Extend work from (Puente, Käser, et al. 2007)

Necessary changes:

- Jacobian matrices  $A$ ,  $B$  and  $C$  more densely populated.
- Flux solver needs an eigendecomposition: switch from analytic expression to numerical solver
- Wave speeds depend on the direction: switch from single evaluation to sampling
- Free surface boundary condition: Solve an inverse Riemann Problem

# Anisotropy – application example



**Figure:** Vertical velocity field after 3 s, left: isotropic material, right: anisotropic material



## Model extensions – Poroelasticity

Poroelastic material: porous elastic medium, filled with a fluid  
In seismology: georeservoirs

Necessary changes:

- $q$  now contains fluid pressure  $p$  and fluid velocities as additional quantities
- Coupling between fluid and solid phase for low frequencies:  
$$\partial_t q + A \partial_x q + B \partial_y q + C \partial_z q = \boxed{Eq}$$
- System of PDEs is stiff!
- ~~reduce time step or~~ use (locally) implicit scheme

## Poroelasticity – Space time predictor

**Before:** Expand the solution from  $t^n$  as Taylor series in time

**Now:**  $q_p(t, \vec{x}) = Q_{pts} \phi_I(\vec{x}) \chi_s(t)$  (Puente, Dumbser, et al. 2008)

- Plug this into the discretization to obtain a linear system of equations.
- For order 6 this has 4368 unknowns.
- We have to solve this system for every element and every timestep.
- Standard approach: precompute LU decomposition, do backsubstitution  
⇒ 38.2 MFLOP for the backsubstitution only

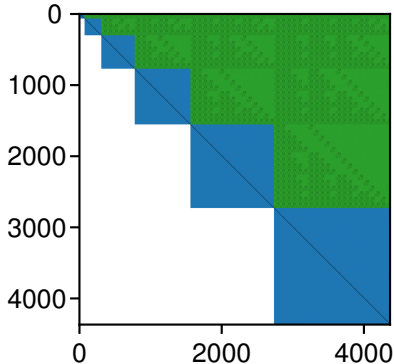
# Poroelasticity – System of equations

Solve a linear equation in tensorial form:  $O_{pksqilt} Q_{qilt} = b_{pks}$ :

$$\begin{aligned}
 & \delta_{pq} \langle \chi_s(\mathbf{1}) \phi_k, \chi_t(\mathbf{1}) \phi_l \rangle Q_{qilt} - \delta_{pq} \left[ \frac{\partial}{\partial \tau} \chi_s \phi_k, \chi_t \phi_l \right] Q_{qilt} \\
 & + \mathbf{A}_{pq}^* \left[ \chi_s \phi_k, \chi_t \frac{\partial}{\partial \xi} \phi_l \right] Q_{qilt} + \mathbf{B}_{pq}^* \left[ \chi_s \phi_k, \chi_s \frac{\partial}{\partial \eta} \phi_l \right] Q_{qilt} \\
 & + \mathbf{C}_{pq}^* \left[ \chi_s \phi_k, \chi_t \frac{\partial}{\partial \zeta} \phi_l \right] Q_{qilt} - \mathbf{E}_{pq}^* [\chi_s \phi_k, \chi_s \phi_l] Q_{qilt} \\
 & = \delta_{pm} \langle \chi_s(\mathbf{0}) \phi_k, \phi_n \rangle Q_{mn}^0.
 \end{aligned}$$

# Poroelasticity – Examine the system more closely

Unroll  $pks$  and  $qIt$  to linear indices:



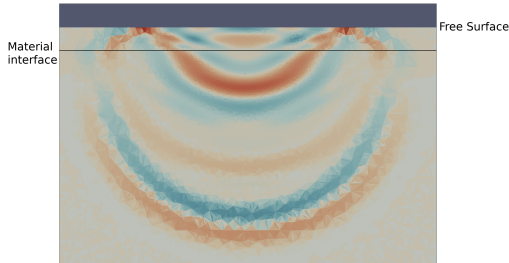
**Figure:** Sparsity pattern of the system matrix, if unrolled correctly

## Poroelasticity – Improved solver

- System is already in upper triangular form, but with blocks of size  $\mathcal{O} \times \mathcal{O}$  on the diagonal.
- Use blockwise backsubstitution to solve the system.  
⇒ 1.94 MFLOP
- This is only 4 % of the original workload.
- Blockwise backsubstitution can be mapped to GEMMs with YATeTo for high performance

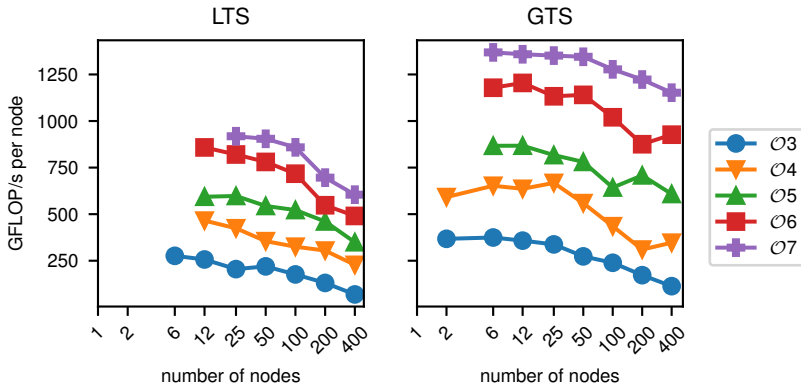
# Poroelasticity – Results

- Convergence test ✓
- Various benchmarks against analytic solutions ✓
- LOHp ✓



**Figure:** Layer over halfspace scenario with poroelastic materials. Vertical velocity after 1 s.

# Poroelasticity – Parallel efficiency



**Figure:** Parallel efficiency for a mesh with 7.33 million elements. We see good scaling until 100.000 elements per node.

# Conclusion

- Successfully added anisotropic material behaviour to SeisSol
- Extension to poroelastic materials is work-in-progress, with promising results



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Algorithm design for the inversion procedure in collaboration with Carsten Uphoff and Martin Galis.

## Backup Slide Equations

Weak formulation of the PDE in 1D:

$$\int_T \partial_t \mathbf{q} \cdot \phi \, dx - \int_T \mathbf{A} \mathbf{q} \partial_x \phi \, dx + \int_{\partial T} \phi \mathbf{A} \mathbf{q} \cdot \mathbf{n} \, ds = \int_T \mathbf{E} \mathbf{q} \phi \, dx$$

Semidiscrete form:

$$\begin{aligned} & \partial_t \mathbf{Q}_{pl} \int_T \phi_l \phi_k \, dx - \mathbf{A}_{pq} \mathbf{Q}_{pl} \int_T \phi_l \partial_x \phi_k \, dx \\ & + \int_{\partial T} \mathbf{F}_{pk}(\mathbf{Q}_{pl}, \mathbf{Q}_{pl}^i) \, ds \\ & = \mathbf{E}_{pq} \mathbf{Q}_{pl} \int_T \phi_l \phi_k \, dx \end{aligned}$$