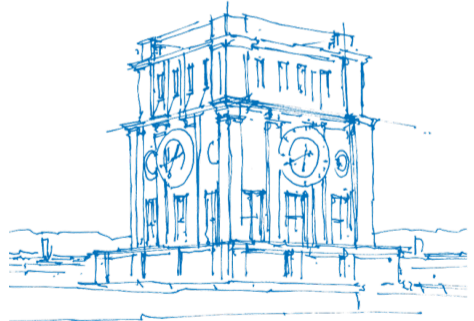


# Advanced Material Models and Dynamic Rupture in SeisSol

SeismoChat, Scripps Institute of Oceanography

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May 26, 2023



*TUM Uhrenturm*

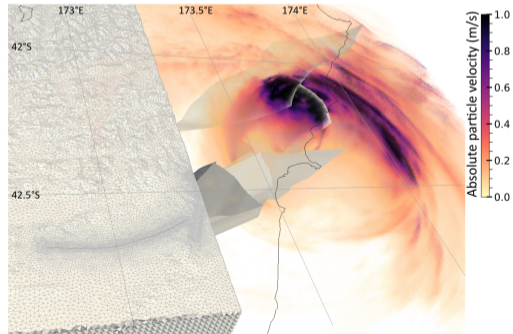
# Earthquake modeling

**Wave propagation:** Solve the elastic wave equation

- $\partial_t q + A\partial_x q + B\partial_y q + C\partial_z q = 0$
- $q$  contains stresses and velocities,  $A$ ,  $B$  and  $C$  contain material information.

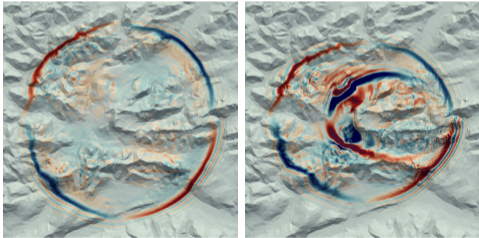
**Source Dynamics:** Prestressed fault and frictional failure criterion

(Ulrich et al., 2019): “Dynamic viability of the 2016 Mw 7.8 Kaikōura earthquake cascade on weak crustal faults”



# Anisotropic materials

- Directional dependent material behaviour, e.g. cracked or layered media
- Jacobian  $A$ ,  $B$ ,  $C$  are more densely populated, but can reuse the numerical scheme from the elastic wave equation (Wolf et al., 2020)



Left: isotropic material, Right:  
anisotropic material

Isotropic:

$$\sigma = \lambda \text{tr}(\epsilon) I + 2\mu \epsilon$$

Anisotropic:

$$\sigma_{ij} = \sum_{k,l=1}^3 c_{ijkl} \epsilon_{kl}$$

# Poroelastic wave equation

- Elastic waves (velocity–stress formulation)
- Fluid phase (pressure and relative fluid velocities)
- Possible applications: Geothermal energy production, fault zone effects
- Challenge: Stiff source term

- Constitutive relation relates stress, strain, pore pressure and fluid variation

$$\begin{pmatrix} \sigma_{xx} \\ \sigma_{yy} \\ \sigma_{zz} \\ \sigma_{yz} \\ \sigma_{xz} \\ \sigma_{xy} \\ -p \end{pmatrix} = \begin{pmatrix} \lambda_M + 2\mu_M + M\alpha^2 & \lambda + M\alpha^2 & \lambda + M\alpha^2 & 0 & 0 & 0 & M\alpha \\ \lambda + M\alpha^2 & \lambda_M + 2\mu_M + M\alpha^2 & \lambda + M\alpha^2 & 0 & 0 & 0 & M\alpha \\ \lambda + M\alpha^2 & \lambda + M\alpha^2 & \lambda_M + 2\mu_M + M\alpha^2 & 0 & 0 & 0 & M\alpha \\ 0 & 0 & 0 & \mu & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & \mu & 0 & 0 \\ 0 & 0 & 0 & 0 & 0 & \mu & 0 \\ M\alpha & M\alpha & M\alpha & 0 & 0 & 0 & M \end{pmatrix} \begin{pmatrix} \epsilon_{xx}^M \\ \epsilon_{yy}^M \\ \epsilon_{zz}^M \\ \epsilon_{yz}^M \\ \epsilon_{xz}^M \\ \epsilon_{xy}^M \\ -\zeta \end{pmatrix}$$

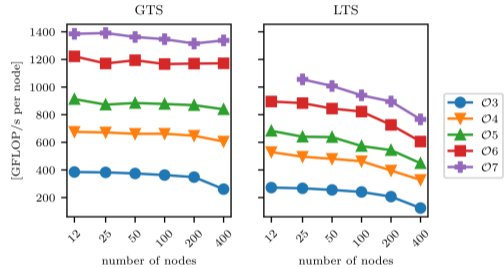
- Algebraic source term couples fluid and solid movement

$$\partial_t \mathbf{q} + \mathbf{A} \partial_x \mathbf{q} + \mathbf{B} \partial_y \mathbf{q} + \mathbf{C} \partial_z \mathbf{q} = \underbrace{\mathbf{E} \mathbf{q}}$$

# HPC aspects

- Global time stepping (GTS)
  - Every element has the same time step
  - regular update scheme
- Local time stepping (LTS)
  - Every element can have its own time step
  - complicated scheduling, but less computational work

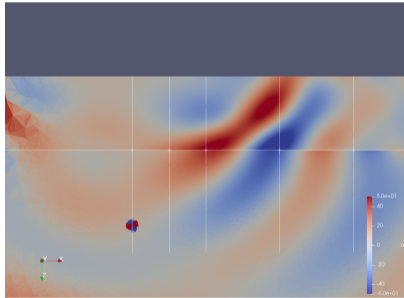
⇒ LTS reduces time-to-solution, but parallel efficiency decreases



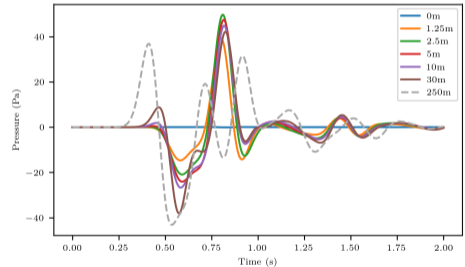
Parallel efficiency of a poroelastic setup with 7.3 million elements for global (GTS) and local (LTS) time-stepping on SuperMUC-NG.

# Challenge: rapidly changing pressure at material interfaces

Layer over halfspace with double-couple point source



Pressure field in the  $x - z$  plane

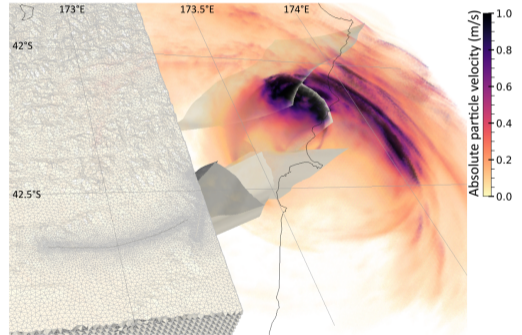


Pressure time history at  $x = 0$ .

# Dynamic Rupture

Internal boundary condition as earthquake source:

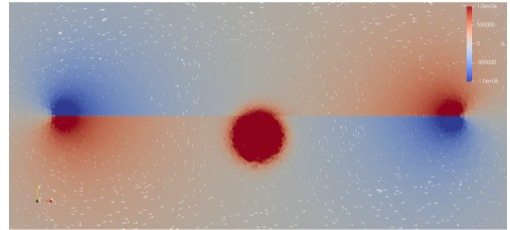
- Compute normal stress  $\sigma_n$  and traction  $\tau$  at the fault.
- Evaluate fault strength  $\mu\sigma_n$ .
- If  $\tau > \mu\sigma$ : evaluate slip rate.
- Impose discontinuous velocity field at the interface



Kaikoura earthquake (Ulrich et al., 2019)

# Poroeelastic Dynamic Rupture

- Internal boundary condition: slip rate based normal stress and tangential traction.
- Pore pressure weakens fault:  
 $\hat{P}_n = \sigma_n - P_f$ .
- Pressure perturbation as initial condition.



Pressure during rupture

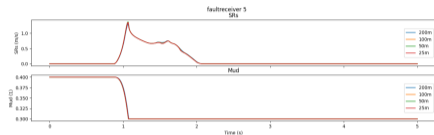
## Project at Scripps

- Convergence analysis
- Parameter study: How does the biot coefficient/porosity change the rupture
- Branched faults

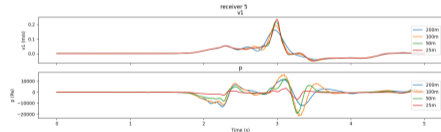


# Convergence analysis

- Homogeneous halfspace, vertical fault (4 km wide)
- Rupture nucleation by fluid stress perturbation (5 MPa,  $r \approx 600$  m)
- Record wave field for 5 s
- Fault: almost perfect results, wave field: could be better



Fault receiver at (1200, 0, -700)



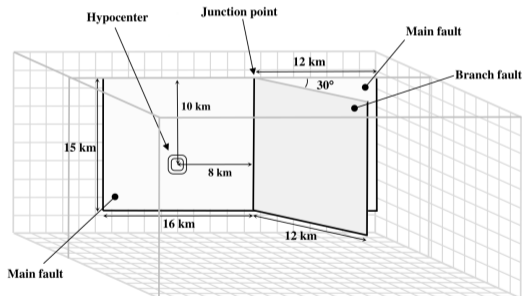
Surface receiver at (-3000, 3000, 0)

# How do poroelastic effects influence branching?

## Geometry

- vertical fault (28 km × 15 km)
- branch at 16 km from the *left* end, at 15°
- nucleation at 8 km from the *left* end, at 10 km depth
- prestress: right lateral rupture:

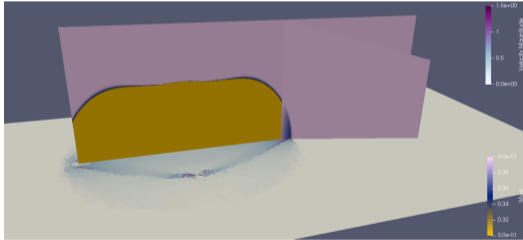
$$\sigma = \begin{pmatrix} -29.5 & 10.0 & 0.0 \\ 10.0 & -25.5 & 0.0 \\ 0.0 & 0.0 & -25.5 \end{pmatrix} \text{ MPa}$$



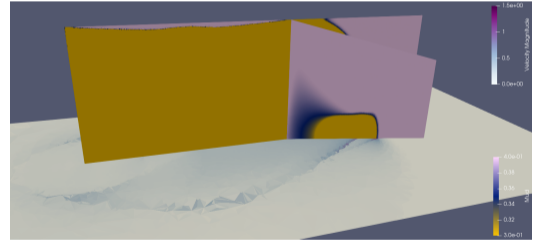
TPV24 geometry with 30° branch angle

# What happens?

Use a material with Biot's coefficient  $\alpha = 1 - \frac{K_M}{K_S} = 0.95$ .



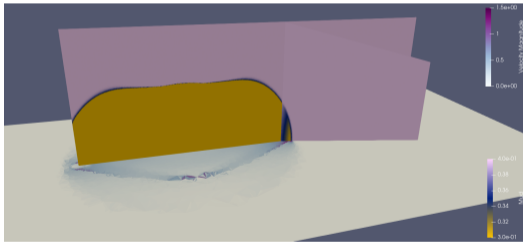
Friction coefficient and wave field after  
2.3 s



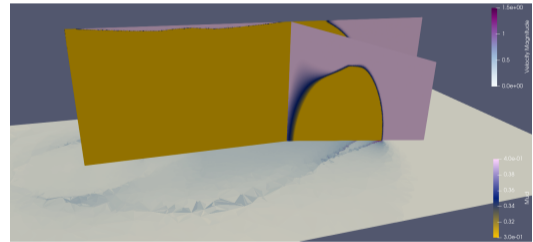
Friction coefficient and wave field after  
4.5 s

## How's the elastic case?

Use the *Gassmann* equivalent elastic material

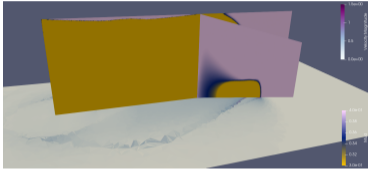


Friction coefficient and wave field after  
2.3 s

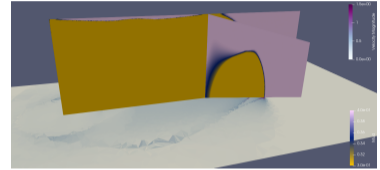


Friction coefficient and wave field after  
4.5 s

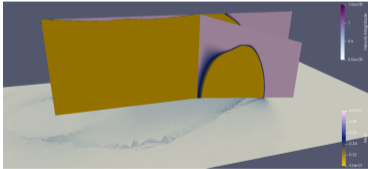
# Parameter study



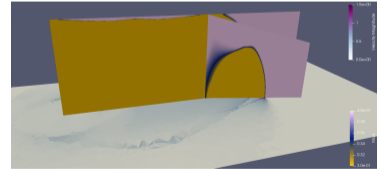
Large poroelastic effect  $\alpha = 0.95$



Small poroelastic effect  $\alpha = 0.1$



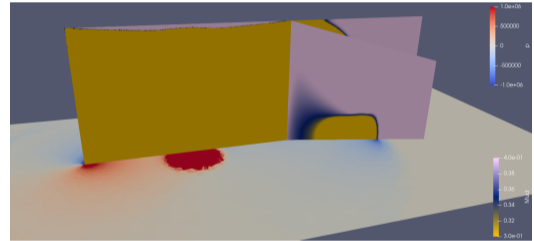
Elastic equivalent



Elastic equivalent

# Why?

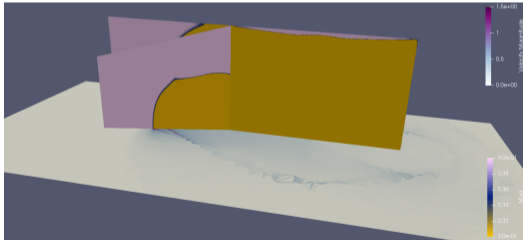
- Branch in the extensive quadrant of the main fault
- $\Delta P < 0 \Rightarrow$  increases  $\hat{P}_n$
- On the other side ( $\phi = -15^\circ$ ), prestress on branch is stronger
- $\Rightarrow$  no rupture on pressure weakening side.



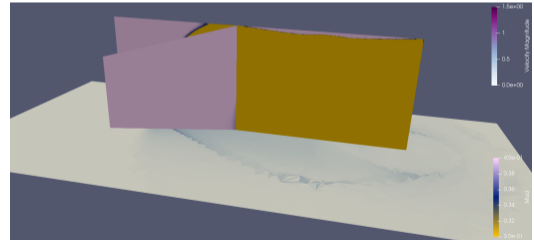
Poroelastic material ( $\alpha = 0.95$ ) with with pore pressure.

# Is there a pressure weakening configuration?

**Yes!** Make  $-15^\circ$  branch weaker:  $\sigma = \begin{pmatrix} -32.8 & 8.7 & 0.0 \\ 8.7 & -22.2 & 0.0 \\ 0.0 & 0.0 & -25.5 \end{pmatrix}$  MPa



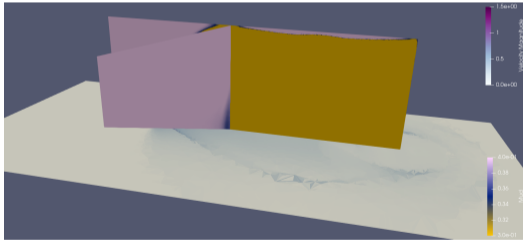
Pressure weakening breaks fault



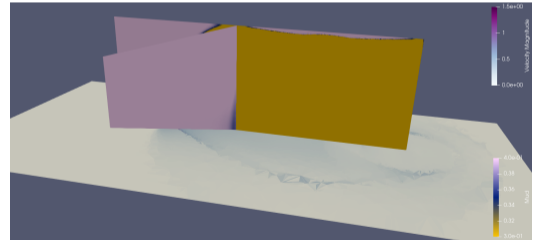
Elastic equivalent doesn't break

# Does this change with $\alpha$ ?

**Yes!**, consider  $\alpha = 0.1$ : the branch never breaks



Pressure weakening doesn't break



Elastic equivalent doesn't break



**Poroelastic effects can inhibit or facilitate fault branching!**

## Take home message

- HPC enables seismological research, but it is hard.
- Domain scientists and computer scientists have to work closely together.
- SeisSol supports complicated material models, where analytic possibilities end.
- Dynamic rupture in poroelastic materials is possible in 3D!
- Pore pressure variation can inhibit or facilitate fault branching

Try SeisSol: <https://github.com/SeisSol/Training>

Ask me: [wolf.sebastian@cit.tum.de](mailto:wolf.sebastian@cit.tum.de) or in the barnyard until June 2.