

3D Acoustic-Elastic Simulations for Tsunami-Genesis

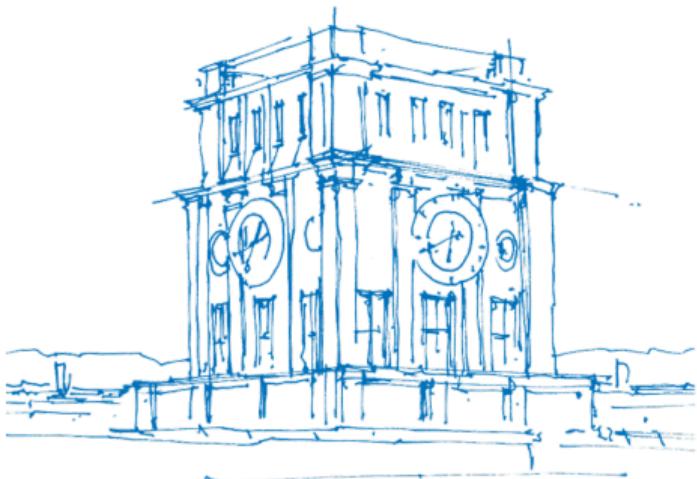
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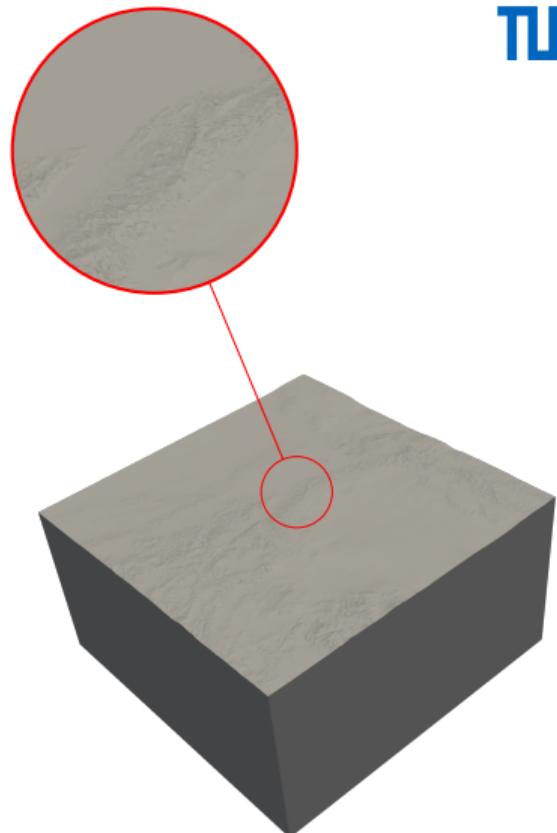
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SeisSol

What

- Seismic wave propagation in (an)**isotropic elastic** material
- **Acoustic** wave propagation
- Viscoelastic wave propagation
- **Off-fault plasticity**
- **Dynamic earthquake rupture**



How

- Numerics: ADER-DG
- Unstructured tetrahedral meshes with local time-stepping
- Optimized Hybrid MPI + OpenMP Parallelization

Palu mesh, 22 mio. elements

Available (**open-source**) at <https://github.com/SeisSol/SeisSol/>.

The ADER-DG Approach¹

Solve linear hyperbolic equations of the form

$$\frac{\partial \mathbf{q}}{\partial t} + \mathbf{Q} \frac{\partial \mathbf{q}}{\partial x} + \mathbf{B} \frac{\partial \mathbf{q}}{\partial y} + \mathbf{C} \frac{\partial \mathbf{q}}{\partial z} = 0 \quad (1)$$

with \mathbf{q} vector of variables, $\mathbf{x} = (x, y, z)$ position, t time, $\mathbf{A}(\mathbf{x})$, $\mathbf{B}(\mathbf{x})$, $\mathbf{C}(\mathbf{x})$ flux matrices.

Discontinuous Galerkin (DG) divides domain into disjoint elements, approximates solutions by piecewise-polynomials.

Elements are connected by solving the Riemann problem.

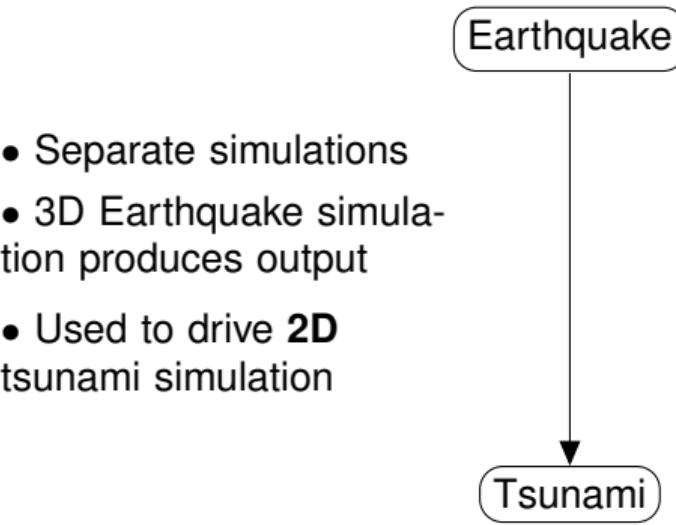
ADER-Approach uses element-local Taylor expansion for time integration instead of Runge-Kutta procedures.

Advantages: One-step scheme, arbitrary order in time and space

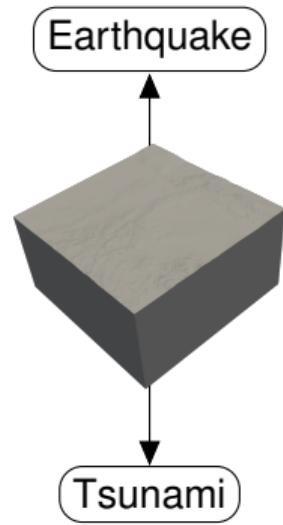
¹V. A. Titarev and E. F. Toro. "ADER: Arbitrary High Order Godunov Approach". In: *Journal of Scientific Computing* 17.1 (Dec. 2002).

Earthquake-Tsunami Coupling Workflows

- ## Two-step Coupling
- Separate simulations
 - 3D Earthquake simulation produces output
 - Used to drive **2D** tsunami simulation



- ## Fully Coupled
- One **3D** simulation



Fully-Coupled

Model

1. From earthquake model: record final seafloor displacement.
2. In tsunami model: Use as initial condition

Model valid when

- In shallow water limit
- Acoustic waves are not generated
- Tsunami waves do not propagate over source duration

Time-dependent Seafloor Velocity

1. From earthquake model: record time-dependent sea floor velocity
2. In tsunami model: use as forcing term

Time-dependent Sea Surface Velocity

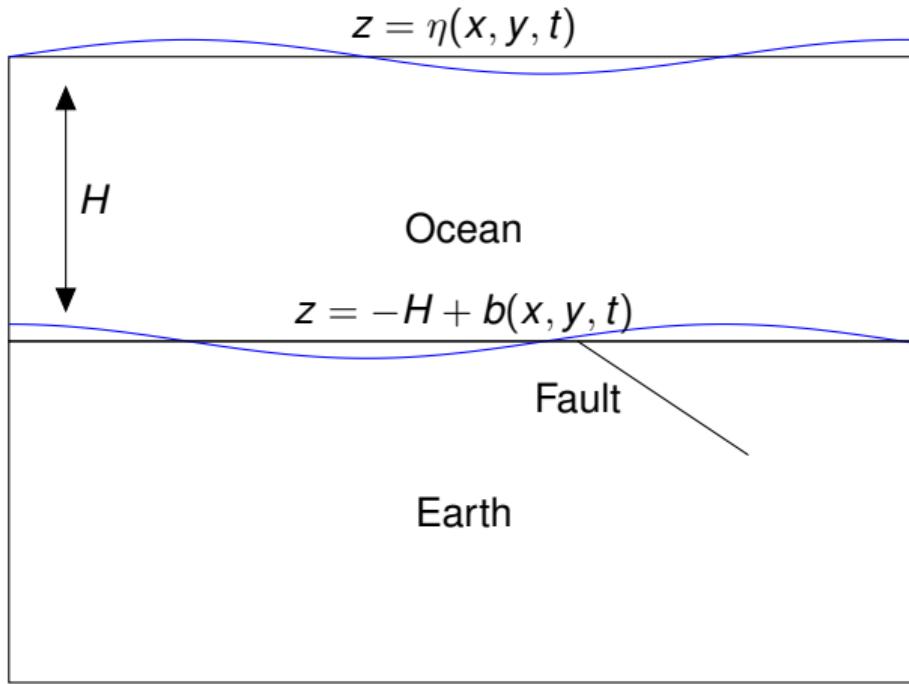
1. From earthquake model: record time-dependent sea surface velocity
2. In tsunami model: use as forcing term

Simultaneously solve earthquake dynamics and ocean response (including gravity)

Valid in all cases

Modified from L. S. Abrahams et al. "Comparison of techniques for coupled earthquake and tsunami modeling". In: AGU Fall Meeting Abstracts (2020)

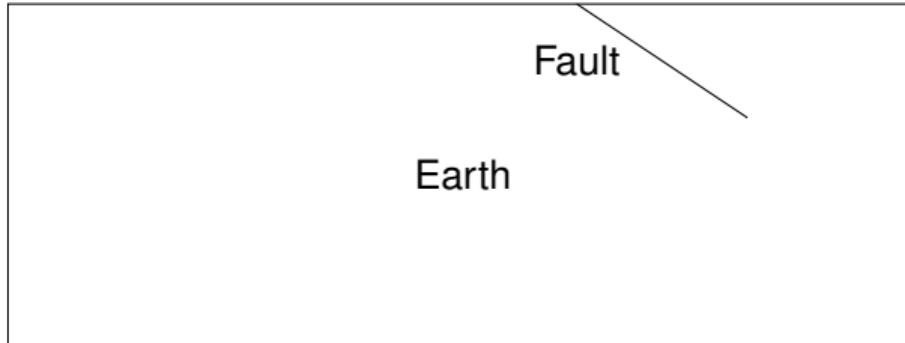
Two-Way Elastic-Acoustic Coupling



With:

- $\eta(x, y, t)$ sea surface height
- H height of the ocean
- $b(x, y, t)$ time dependent bathymetry
- Figures in 2D for illustration, all simulations are **3D**

Earth Model



$$\frac{\partial \sigma_{ij}}{\partial t} - \lambda \delta_{ij} u_k \frac{\partial}{\partial x_k} - \mu \left(\frac{\partial}{\partial x_j} u_i + \frac{\partial}{\partial x_i} u_j \right) = 0, \quad (2)$$
$$\rho \frac{\partial}{\partial t} u_i - \frac{\partial}{\partial x_j} \sigma_{ij} = 0$$

δ_{ij} Kronecker delta, summation implied

- Isotropic elastic medium
- Velocity-stress formulation
- u, v, w velocities
- σ stress tensor
- ρ density, (μ, λ) Lame parameters
- Dynamic rupture earthquake source (e.g. rate & state friction)

Ocean Model²

Modeled as **linear acoustic** medium, $\mathbf{q} = (u, v, w, p)$

Treated as special case of elastic wave equation with $\mu = 0$.

Pressure p sum of background pressure p_0 (in **hydrostatic equilibrium**) and perturbation p' .

$$\begin{aligned} p &= p_0 + p'(x, y, z) \\ p_0 &= p_a + \rho g(-z) \end{aligned} \tag{3}$$

with atmospheric pressure p_a and $g = 9.81 \text{ m/s}^2$.

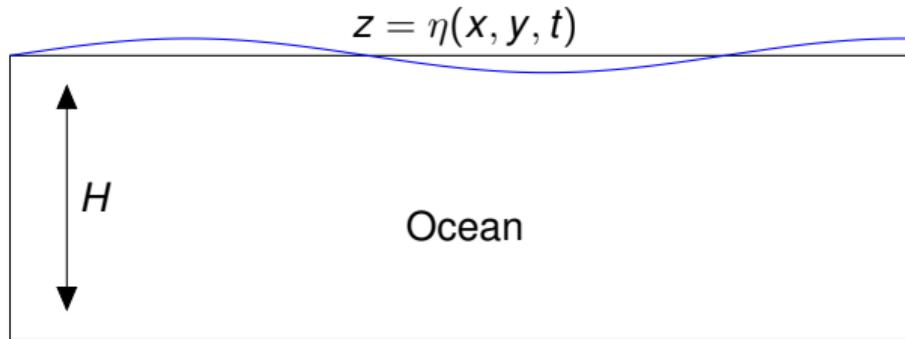
Pressure at some point is:

$$p(x, y, z, t) = p_a + \rho g(-z) + p'(x, y, z, t) - \rho g u_z(x, y, z, t) \tag{4}$$

with z -displacements u_z .

²G. C. Lotto and E. M. Dunham. "High-order finite difference modeling of tsunami generation in a compressible ocean from offshore earthquakes". In: Computational Geosciences 19.2 (2015).

Ocean Model: Free Surface



Physical free surface boundary condition at sea surface height η :

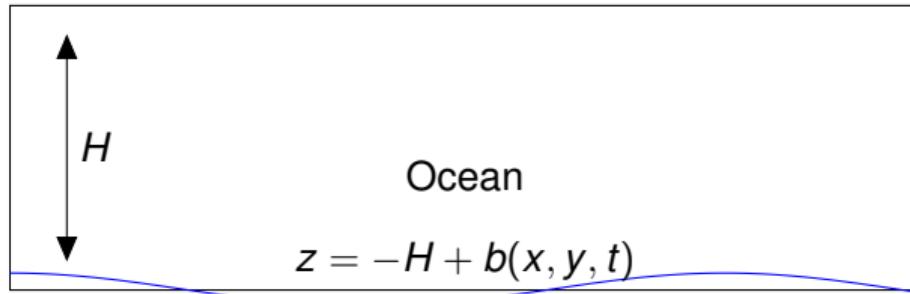
$$p(x, y, \eta) = 0 \tag{5}$$

Too expensive, instead use **linearization** and **hydrostatic assumption**:

$$\begin{aligned} p(x, y, z = 0) &= -\rho g \eta(x, y) \\ \frac{\partial \eta}{\partial t} &= u \end{aligned} \tag{6}$$

Important to use u at boundary (solution of Riemann problem), otherwise unstable!

Ocean Model: Sea floor



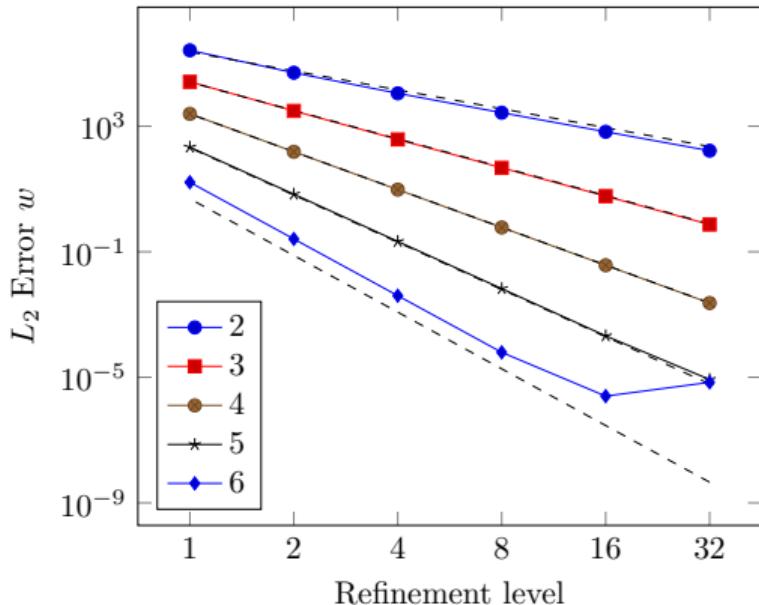
For sloping $b(x, t)$ we have:

$$b(x, y, t) \approx u_z(x, y, t) - \frac{dH}{dx} u_x(x, y, t) - \frac{dH}{dy} u_y(x, y, t), \quad (7)$$

Our method contains typical Tanioka & Satake³ correction!

³Y. Tanioka and K. Satake. "Tsunami generation by horizontal displacement of ocean bottom". In: *Geophysical research letters* 23.8 (1996).

Convergence⁴



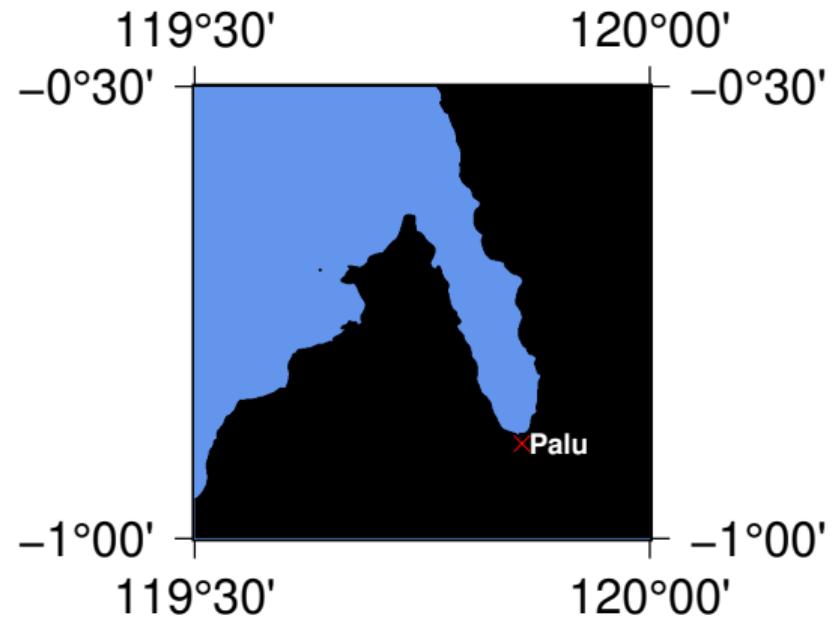
Model verification scenario:

- Acoustic only model
- Simple water tank
- Corresponds to acoustic wave mode in ocean
- Boundary conditions: Modified free surface top, bottom rigid, other free surface
- Simulation time 2.562 s, one wave period
- Number of elements from 600 to ca. 19.7 million elements

⁴L. S. Abrahams et al. "Verification of a 3D Fully-Coupled Earthquake and Tsunami Model". In: AGU Fall Meeting Abstracts 43 (Dec. 2019).

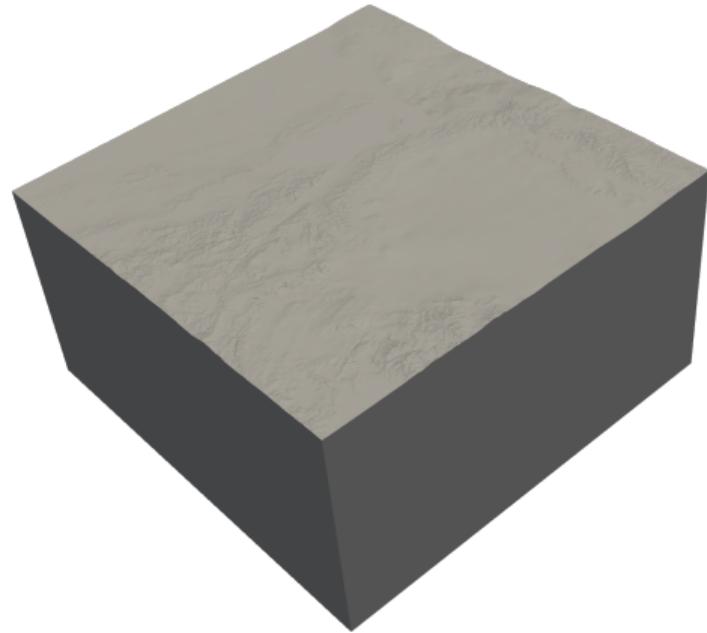
Example: Palu, Sulawesi September 2018

- $M_w 7.5$ strike-slip earthquake
- Propagation at supershear speed crossing narrow Palu bay
- Followed by unexpectedly localized tsunami
- Complicated geometry (bath-tub like bay, very shallow water)



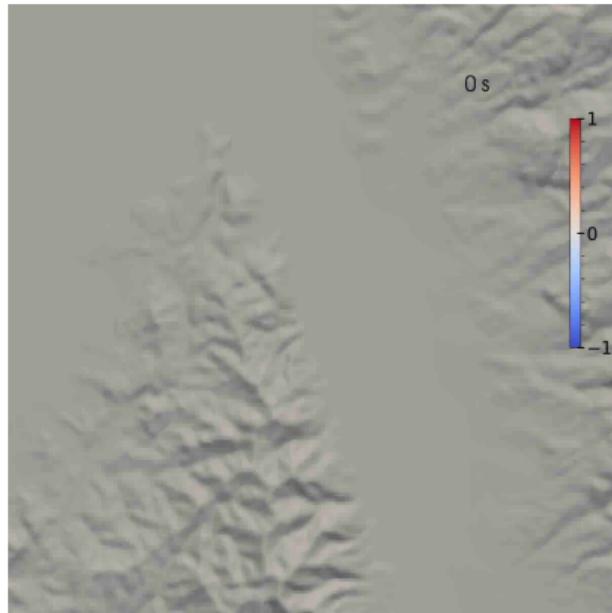
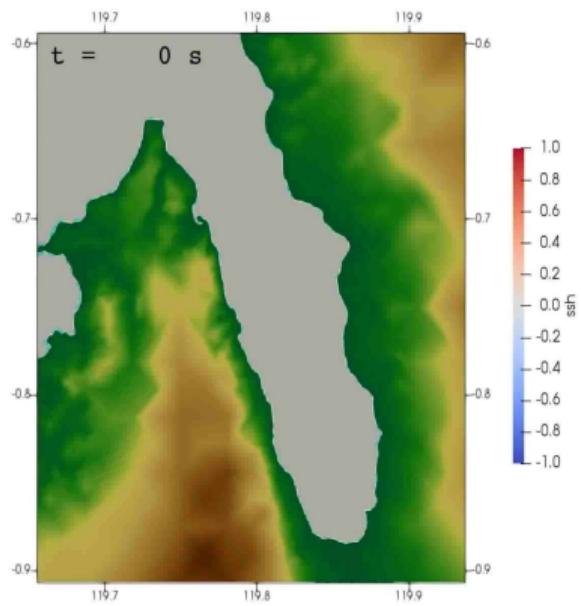
Our setup

- Added water layer to existing earthquake model⁵.
- Fully coupled model (including plasticity, dynamic earthquake rupture)
- Resulting mesh roughly **22 million elements**, used order 6
- Simulation time: 150 s
- Took 6h28min on **256 nodes** of SuperMUC-NG (2 Intel Skylake Xeon Platinium 8174 CPUs)



⁵T. Ulrich et al. "Coupled, Physics-based Modelling Reveals Earthquake Displacements are Critical to the 2018 Palu, Sulawesi Tsunami". In: *Pure and Applied Geophysics* (2019)

Results



Left: Two-step coupling, right: fully-coupled

Conclusion

- Fully coupled elastic-acoustic simulations capture more effects than typical two-step strategies
- Linearization of free surface boundary conditions efficient way of tracking sea surface height
- Our method achieves high-order convergence
- Results for Palu scenario are very promising
- Further work on numerics, performance and scenarios