



Regional Gravity field modeling as multi-resolution representation estimated from the combination of heterogeneous data sets

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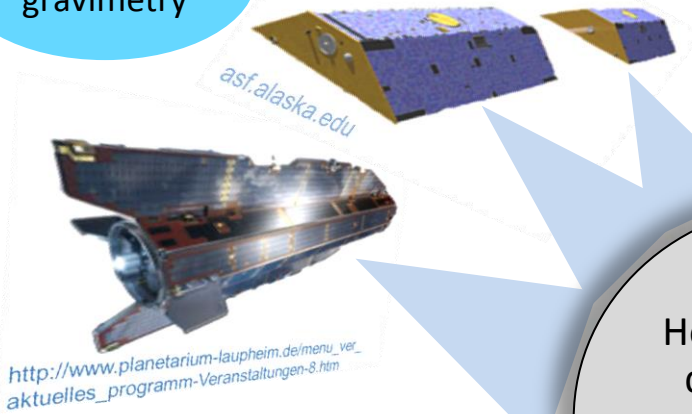
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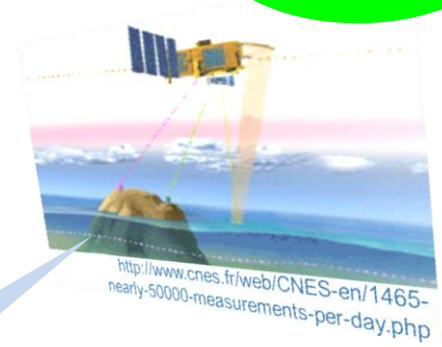
Motivation

Task: Software development for the flexible computation of regional gravity and geoid models from ...

Satellite gravimetry



Altimetry



Difficulty:
How to combine data sets with different distribution, resolution, and accuracy?

airborne gravimetry



terrestrial gravimetry



<http://geogramamente.wordpress.com/2009/03/19/1621/>

... heterogeneous data sets.

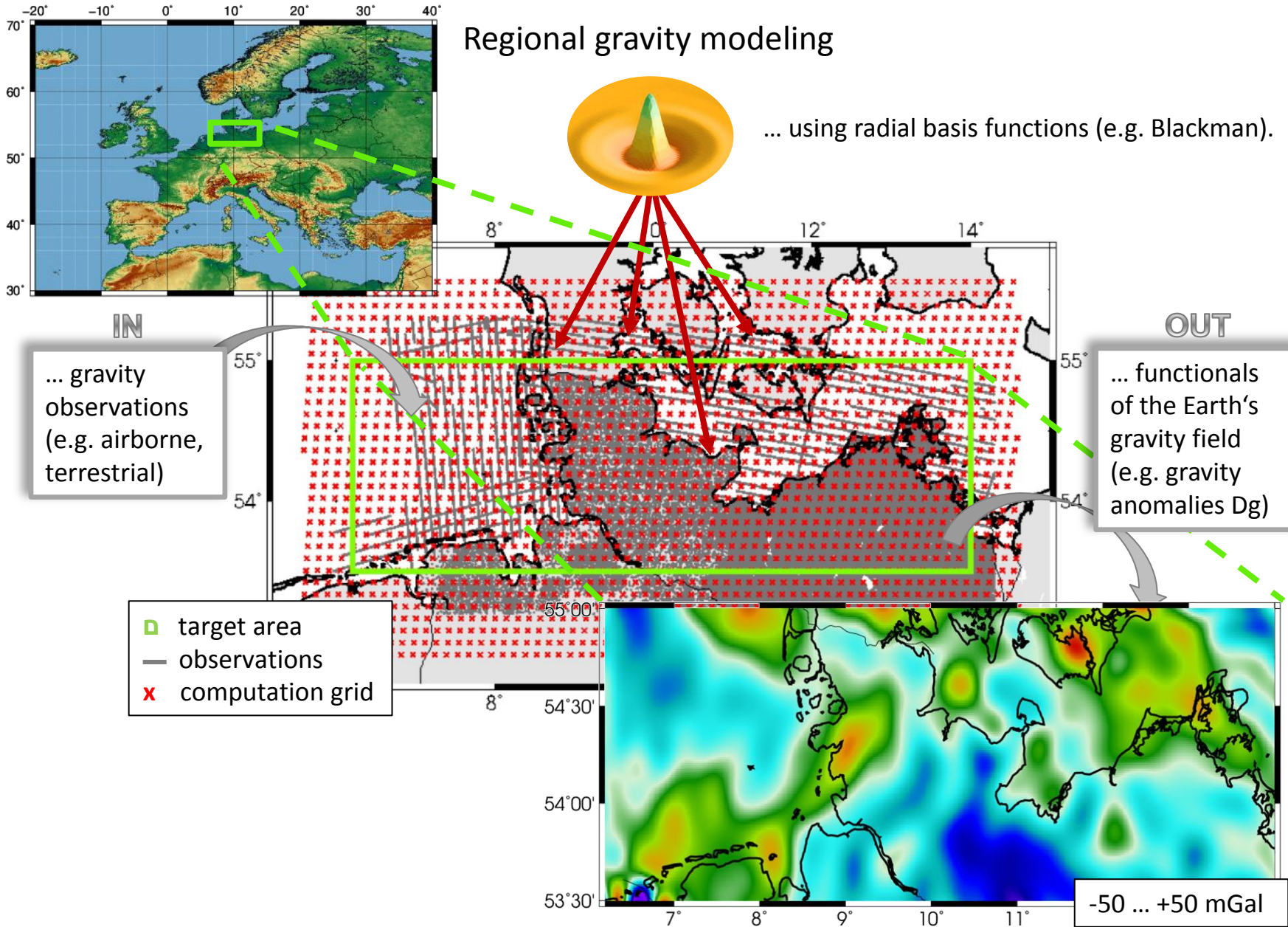
Aim: Construction of precise height systems in operational areas.



Approach

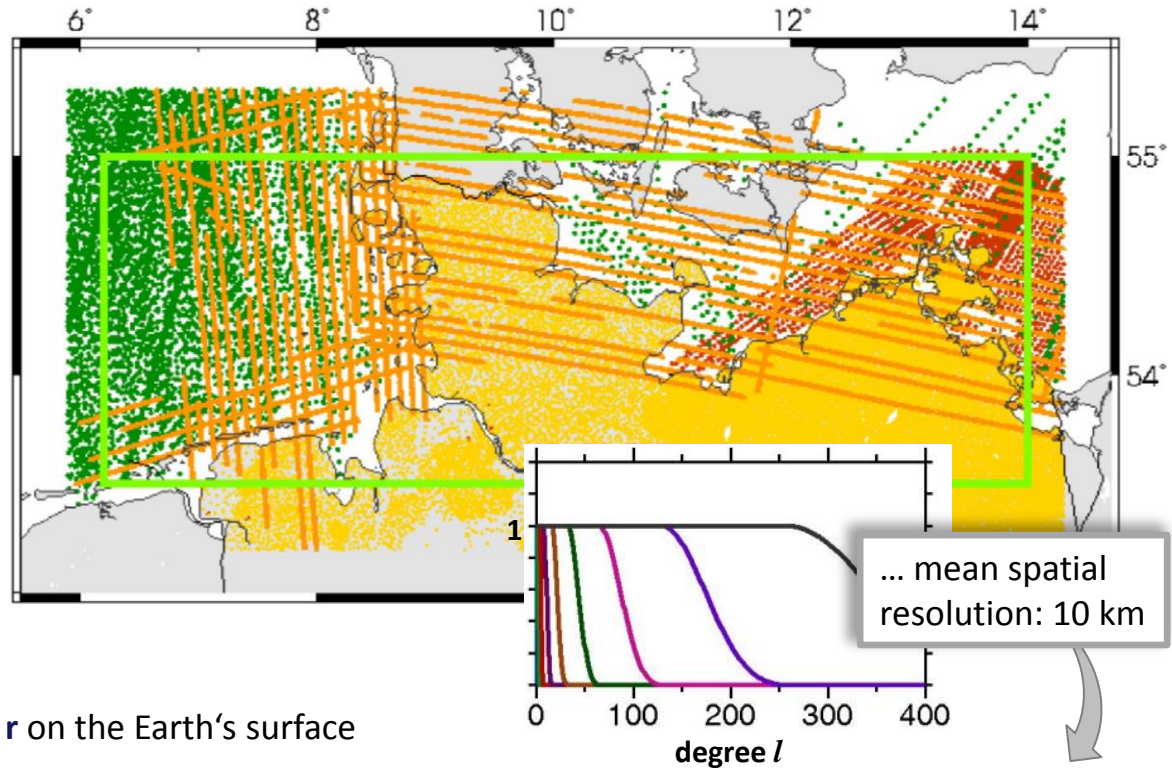
Regional gravity modeling

... using radial basis functions (e.g. Blackman).



Observations

- GOCE GG
- altimetry,
- airborne,
- sea ground,
- terrestrial measurements



... sensitive to different frequency bands
 ... defined by **resolution levels j**
 ... limited by **maximum degree L_j** in a series expansion
 ... related to the **spatial resolution r** on the Earth's surface

j [level]	1	2	3	4	5	6	7	8	9	10	11	...
L_j [deg]	1	3	7	15	31	63	127	255	511	1023	2047	...
r [km]	20000	6667	2857	1333	645	317	157	78	39	20	10	...
		satellite gravimetry										
								altimetry				
							airborne + terrestrial gravimetry					

frequency \rightarrow



Combination of data sets

$$\Delta \mathcal{F}(\mathbf{x}) = \sum_{q=1}^Q d_{J,q} b_J(\mathbf{x}, \mathbf{x}_q) = \sum_{q=1}^Q d_{J,q} \sum_{l=0}^{L_J} \frac{2l+1}{4\pi} \left(\frac{R}{r}\right)^{l+1} B_l P_l(\cos \theta)$$

GOCO03s¹ up to d/o 127

¹Combination of GRACE, GOCE, SLR, ...

technique

Modified basis function \tilde{b}_J

GRACE

ΔV

$$\tilde{b}(\mathbf{x}_1, \mathbf{x}_2, \mathbf{x}_q) = \sum_{l=0}^L \frac{2l+1}{4\pi} \left[\left(\frac{R}{r_1}\right)^{l+1} B_l P_l(\cos \theta_1) - \left(\frac{R}{r_2}\right)^{l+1} B_l P_l(\cos \theta_2) \right]$$

GOCE

V_{zz} (e.g.)

$$\tilde{b}(\mathbf{x}, \mathbf{x}_q) = \sum_{l=0}^L \frac{2l+1}{4\pi} \frac{(l+1)(l+2)}{r^2} \left(\frac{R}{r}\right)^{l+1} B_l P_l(\cos \theta)$$

Altimetry

$N = \text{SSH} - \text{DOT}$

$$\tilde{b}(\mathbf{x}, \mathbf{x}_q) = \sum_{l=0}^L \frac{2l+1}{4\pi} \frac{1}{\gamma} \left(\frac{R}{r}\right)^{l+1} B_l P_l(\cos \theta)$$

Terrestrial, air-,
shipborne,
bathymetry
measurements

Dg

$$\tilde{b}(\mathbf{x}, \mathbf{x}_q) = \sum_{l=0}^L \frac{2l+1}{4\pi} \frac{l-1}{r} \left(\frac{R}{r}\right)^{l+1} B_l P_l(\cos \theta)$$

δg

$$\tilde{b}(\mathbf{x}, \mathbf{x}_q) = \sum_{l=0}^L \frac{2l+1}{4\pi} \frac{1-l}{r} \left(\frac{R}{r}\right)^{l+1} B_l P_l(\cos \theta)$$

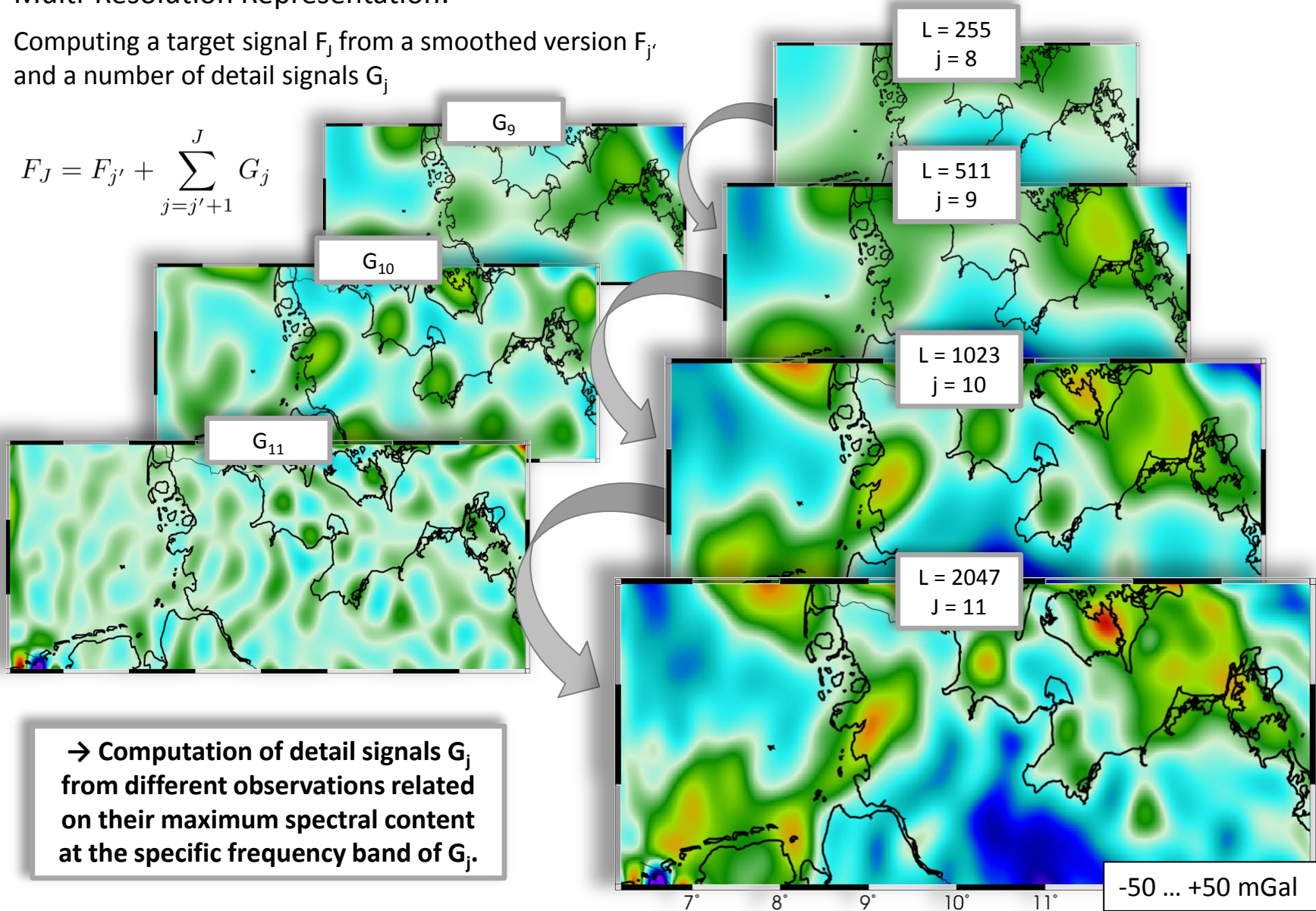
➤ **Estimation of unknown scaling coefficients d_J**

... by using an extended Gauß-Markov model and VCE (rigorous combination at one level j).

Multi-Resolution Representation:

Computing a target signal F_J from a smoothed version $F_{j'}$ and a number of detail signals G_j

$$F_J = F_{j'} + \sum_{j=j'+1}^J G_j$$



→ Computation of detail signals G_j from different observations related on their maximum spectral content at the specific frequency band of G_j .

Relative weighting

Observation	$j = 8$ (L = 255)	$j = 9$ (L = 511)	$j = 10$ (L = 1023)	$j = 11$ (L = 2047)
GOCE V_{xx}	1	10^{-1}	10^{-3}	10^{-5}
GOCE V_{xy}	10^{-4}	10^{-5}	10^{-7}	10^{-9}
GOCE V_{xz}	10^{-1}	10^{-2}	10^{-4}	10^{-6}
GOCE V_{yy}	1	10^{-1}	10^{-3}	10^{-5}
GOCE V_{yz}	10^{-4}	10^{-5}	10^{-7}	10^{-9}
GOCE V_{zz}	1	10^{-1}	10^{-3}	10^{-5}
ERS-1e	1	1	10^{-2}	10^{-3}
ERS-1f	1	1	10^{-2}	10^{-3}
Jason 1 GM	1	1	10^{-1}	10^{-3}
Envisat EM	1	1	10^{-1}	10^{-3}
Cryosat RADS	1	1	10^{-2}	10^{-3}
Airb. North Sea	10^{-1}	1	1	1
Airb. Baltic Sea	10^{-1}	1	10^{-1}	10^{-2}
Terrestrial Data	10^{-1}	1	1	1
Bathymetry	10^{-2}	10^{-1}	1	10^{-1}
Prior information GOCO03s d/o 127	10^{-3}	10^{-4}	10^{-4}	10^{-5}

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Prior information GOCO03s d/o 127	10^{-3}	10^{-4}	10^{-4}	10^{-5}

Criteria

● ● ● ● high sensitivity

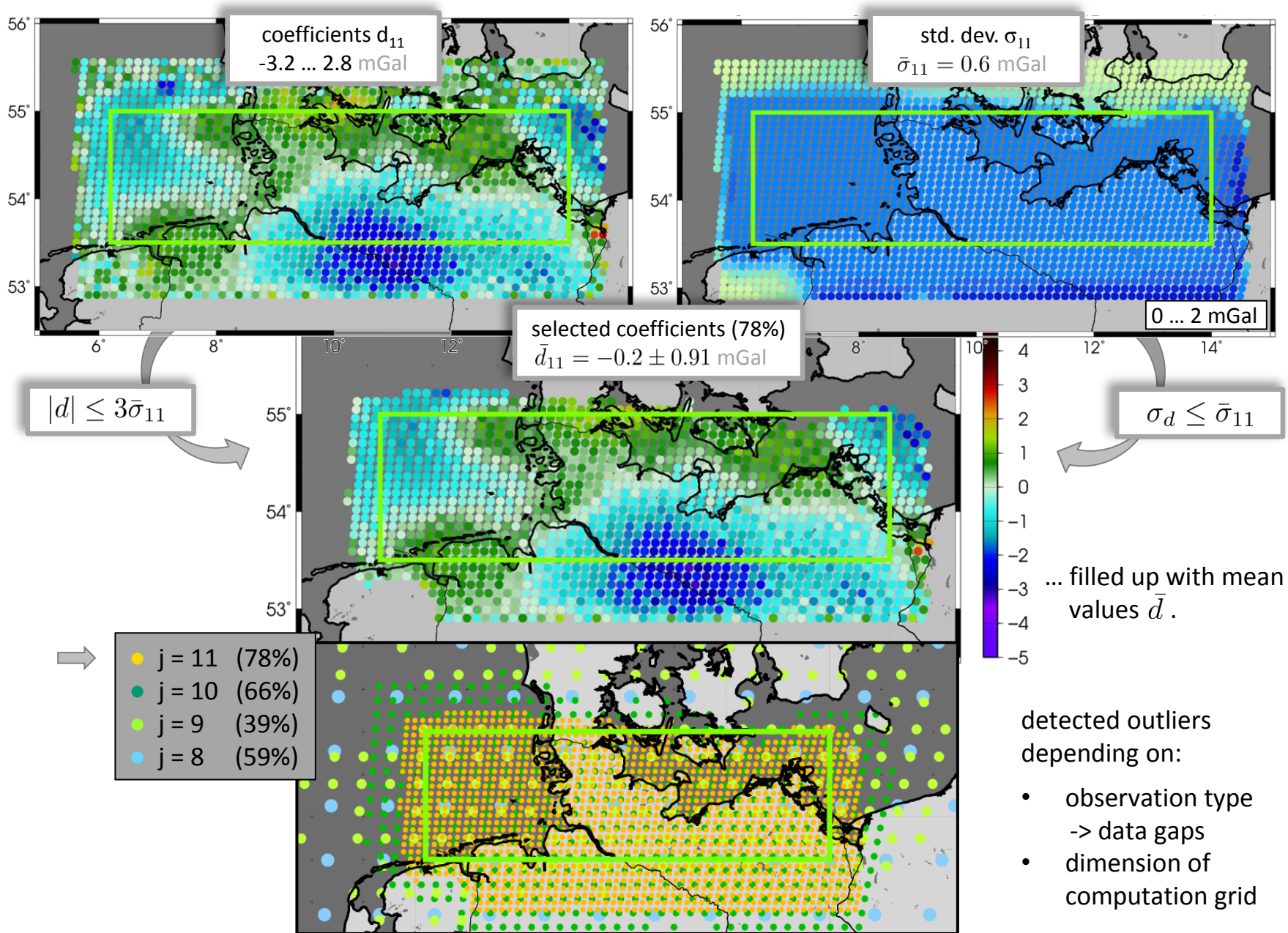
Relative weighting

Observation	j = 8 (L = 255)	j = 9 (L = 511)	j = 10 (L = 1023)	j = 11 (L = 2047)
GOCE V_{xx}	1			
GOCE V_{xy}	10^{-4}			
GOCE V_{xz}	10^{-1}			
GOCE V_{yy}	1			
GOCE V_{yz}	10^{-4}			
GOCE V_{zz}	1			
ERS-1e		1	10^{-2}	10^{-3}
ERS-1f		1	10^{-2}	10^{-3}
Jason 1 GM		1	10^{-1}	10^{-3}
Envisat EM		1	10^{-1}	10^{-3}
Cryosat RADS		1	10^{-2}	10^{-3}
Airb. North Sea			1	1
Airb. Baltic Sea			10^{-1}	10^{-2}
Terrestrial Data			1	1
Bathymetry				10^{-1}
Prior information GOCO03s d/o 127	10^{-3}	10^{-4}	10^{-4}	10^{-5}

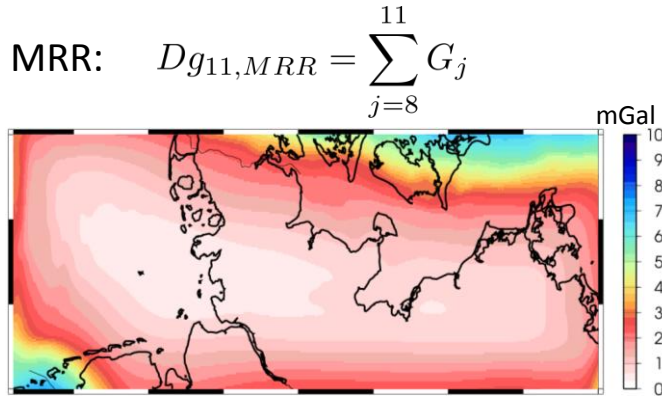
Criteria

- ● ● ● high sensitivity
- ● no correlations
- ● spatial distribution (prior information not sufficient)

Coefficients

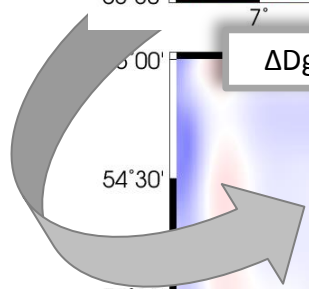
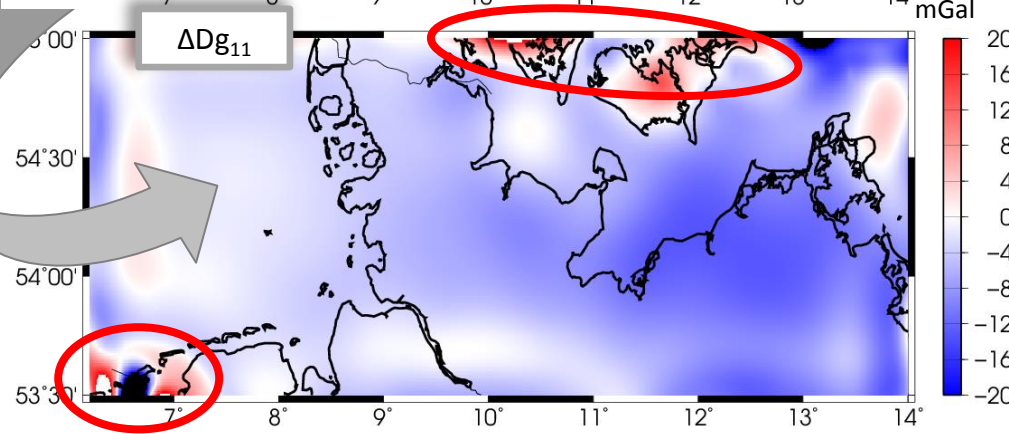
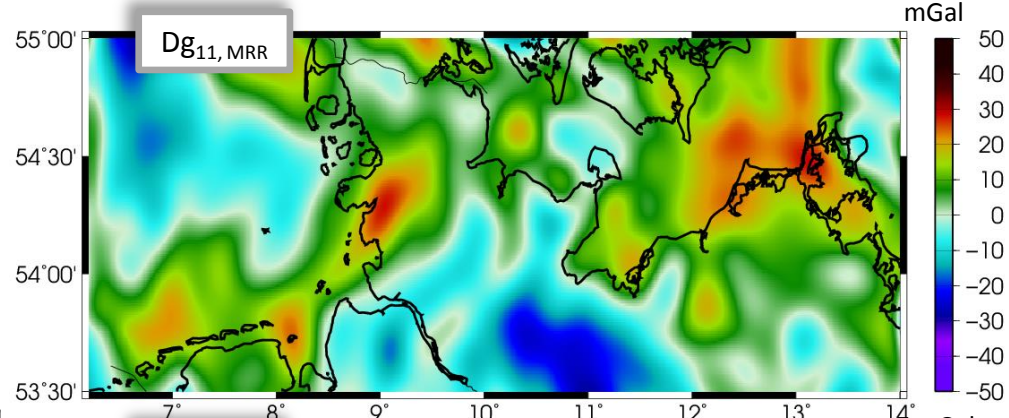
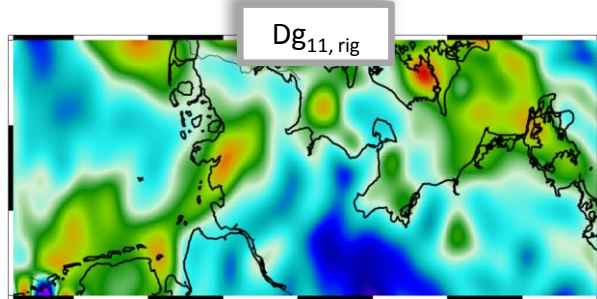


Summation of detail signals



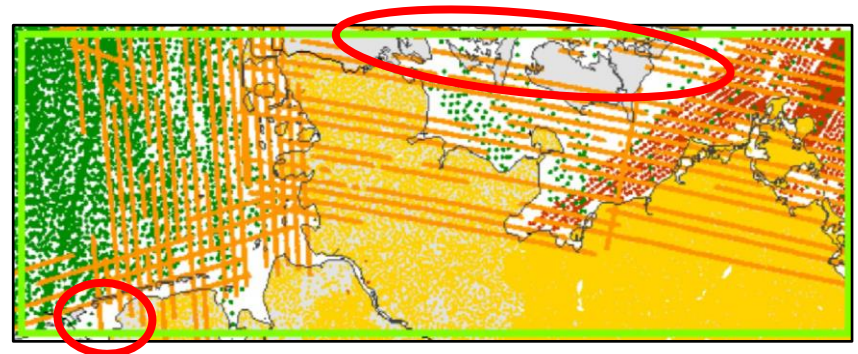
Dg_{11} : -28.41 ... 28.47 mGal
 σ_{11} : 0.46 ... 7.22 mGal

- Largest standard deviations in less observed regions!



ΔDg , mean +/- std: **-4.79 +/- 4.82** mGal

- Largest differences at data gaps close to the borderline.
- **Improvement:** MRR-solution contains optimized spectral information in all frequency domains (contribution of GOCE)!



Outlook

- improving selection of input data
- choosing prior information with higher spectral content (e.g. topographic models)
- considering correlations between detail signals (e.g. introducing a filter matrix)
- improving outlier detection
- validation with real data
- further study areas
- ...

Criteria

- high sensitivity
- no correlations
- spatial distribution (prior information not sufficient)

Summary

Rig. combination @ $j = 11$

- + less unknowns to estimate
- relative weighting of obs. at highest level

MRR combination up to $j = 11$

- larger number of unknowns
- + relative weighting of obs. at each level
- + spectral information in all frequency bands
- + improved handling of data gaps
- stabilized solution

- **Exploiting the highest degree of information out of each data set.**

Appendix

Comparison with EGM2008

$$Dg_{11,final} = GOCO03s + \sum_{j=8}^{11} G_j$$

Difference $Dg_{11, MRR} - EGM2008$

($j = 11, l = 2023$, Blackman smoothed)

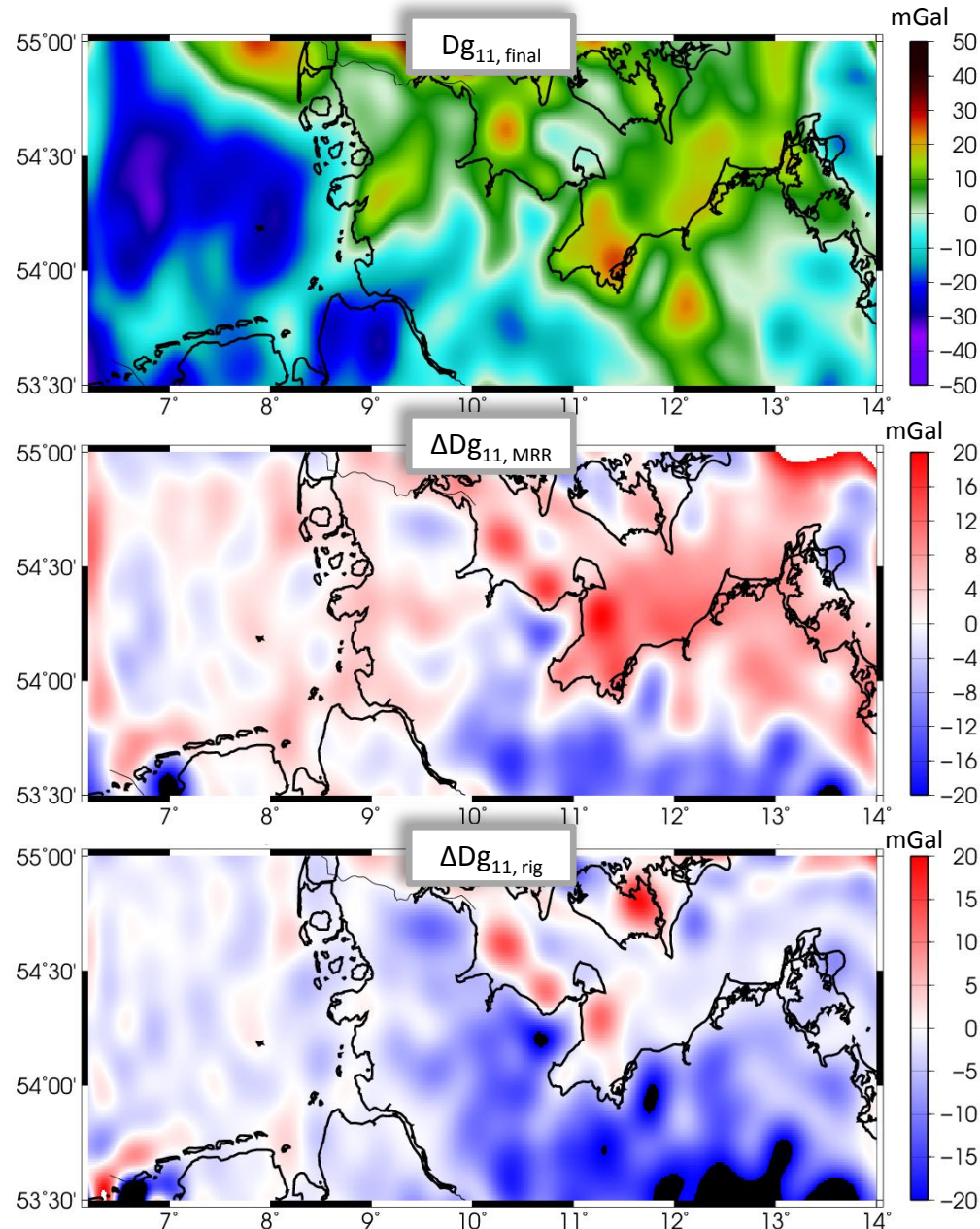
ΔDg , mean +/- std: **0.86 +/- 5.80** mGal

- Largest differences at data gaps.
- Differences up to +/- 10 mGal in western parts (new data set?)
- and in the Baltic Sea (missing airborne data in EGM?).

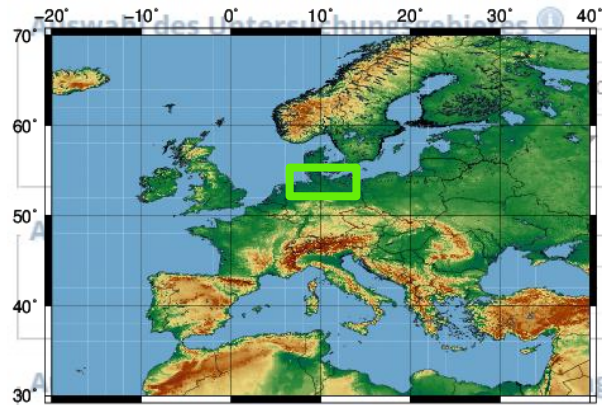
Difference $Dg_{11, rig} - EGM2008$

ΔDg , mean +/- std: **-3.93 +/- 6.01** mGal

- Larger differences (especially in western parts).
- Missing spectral information in mid and low frequency domains.



Software – Specifications

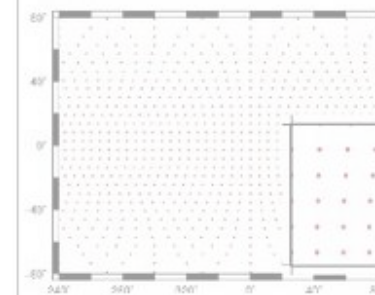


Study area: 6.2° ... 14.0° longitude
53.5° ... 55.0° latitude

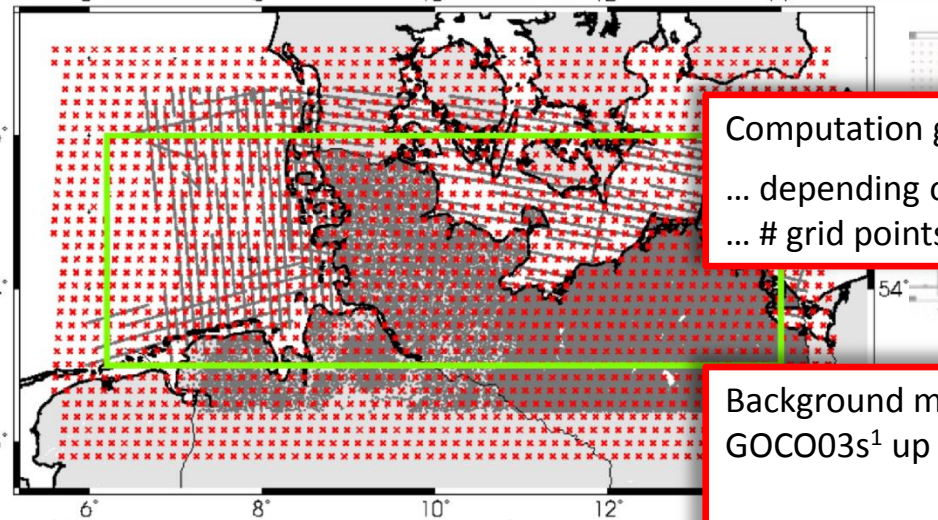
$L_{11} = 2047$ ($J = 11$)

... depending on (spectral/spatial)
resolution of input data

Reuter-Gitter



Zufällige Punktverschiebu



Computation grid: Reuter

... depending on resolution level J
... # grid points = # unknowns Q

Auswahl eines Hintergrund

Schwerefeldmodell: goco03s.250.gfc

Min. Grad: 0

Max. Grad: 127

Background model:
GOCO03s¹ up to d/o 127

... further serves as
prior information due to
rank deficiency problems

¹Combination of GRACE, GOCE, SLR, ...

Auftrag abschicken

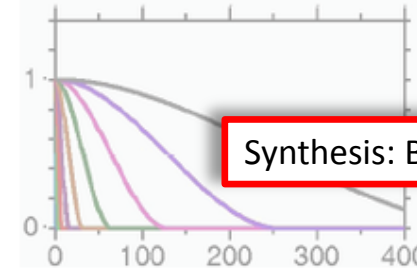
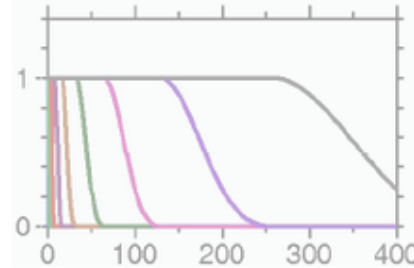
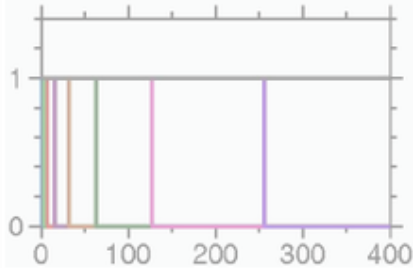
Eingaben löschen

Auswahl der Level ⓘ

Level: 8 - 11

Computation of the target signal
up to max. level $J = 11$

Auswahl der Wavelet- und Skalierungsfunktion ⓘ

 Shannon Blackman Cubic Polynominal*(Darstellung im Frequenzbereich)*

Synthesis: Blackman scaling funct.

Auswahl des Ausgabe-Funktional ⓘ

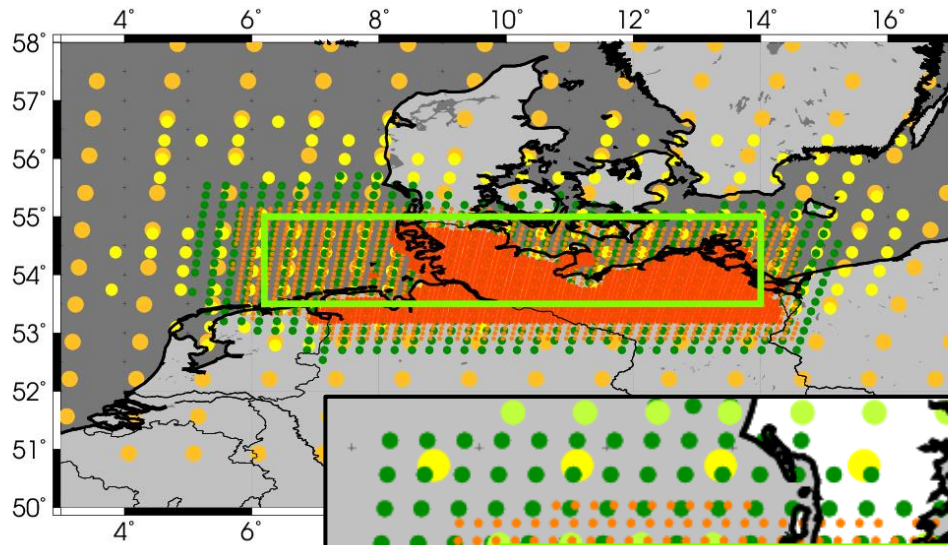
- N** Geoidhöhen [m]
- Dg** Schwereanomalien [mGal]
- V** Potential [m²/s²]
- T** Störpotential [m²/s²]
- Txx** Ableitung $\partial^2 T / \partial x^2$ [E]
- Txy** Ableitung $\partial^2 T / \partial x \partial y$ [E]
- Txz** Ableitung $\partial^2 T / \partial x \partial z$ [E]
- Tyy** Ableitung $\partial^2 T / \partial y^2$ [E]
- Tyz** Ableitung $\partial^2 T / \partial y \partial z$ [E]
- Tzz** Ableitung $\partial^2 T / \partial z^2$ [E]

Output:
... different functionals of the
Earth's gravity field
(e.g. gravity anomalies Dg)

Auswahl des Ausgabegebiets ⓘ

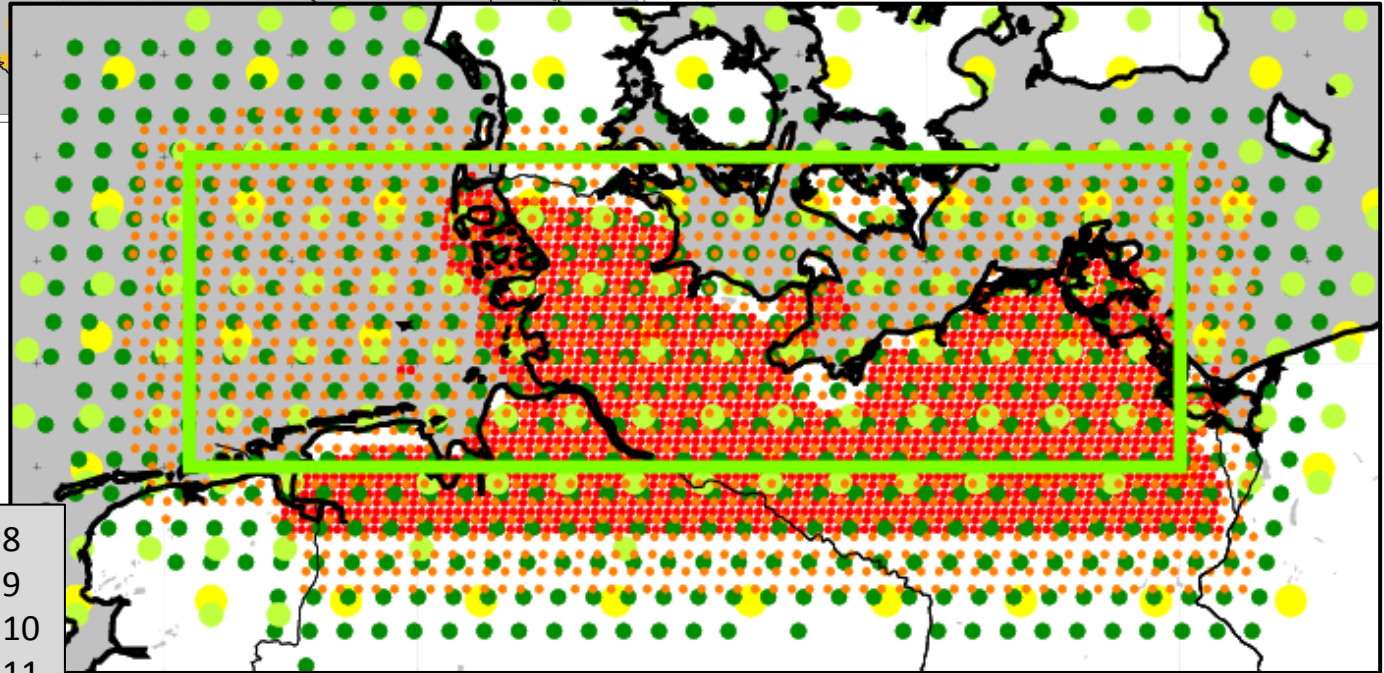
- Gitter aus Gebietsdefinition [km]
- Punkteliste

Coefficients



$$|\sigma_d| \leq \bar{\sigma}_d$$

$$|d| \leq 3\bar{\sigma}_d$$



- $j = 8$
- $j = 9$
- $j = 10$
- $j = 11$
- $j = 12$

$$\Delta \mathcal{F}(\mathbf{x}) = \sum_{q=1}^N d_{J,q} b_{J+1}(\mathbf{x}, \mathbf{x}_q) = \sum_{q=1}^N \sum_{l=0}^{L_J} \frac{2l+1}{4\pi} d_{J,q} \Phi_{J+1,l} \left(\frac{R}{r}\right)^{l+1} P_l(\cos \psi)$$

Observation equation for one observation:

Deterministic part

$$\Delta \mathcal{F}(\mathbf{x}) + e(\mathbf{x}) = \mathbf{b}_{J+1}^T(\mathbf{x}, \mathbf{x}_q) \mathbf{d}_J$$

IN:	$\Delta \mathcal{F}$	observation
	e	measurement error
	\mathbf{b}_{J+1}	(Nx1) vector of basis functions
OUT:	$\hat{\mathbf{d}}_J$	(Nx1) vector of scaling coefficients

Stochastic part

$$D(\Delta \mathcal{F}_k) = \sigma_k^2 \mathbf{P}_k^{-1}$$

IN:	$\Delta \mathcal{F}_k$	vector of observations
	\mathbf{P}_k	weighting matrix of observations
OUT:	$\hat{\sigma}_k$	variance components (VCs)

Estimation of unknown scaling coefficients \mathbf{d}_J

Introduction of additional observations $\boldsymbol{\mu}_d$

$$\boldsymbol{\mu}_d + \mathbf{e}_d = \mathbf{d}$$

with $D(\boldsymbol{\mu}_d) = \sigma_d^2 \mathbf{P}_d^{-1}$

- $\boldsymbol{\mu}_d$... prior information
- avoiding singularity problems
- rank deficiencies (in general number of grid points too large)

Extended Gauß-Markov model for several observation techniques:

$$\mathbf{y} = \begin{bmatrix} \Delta V_1 \\ \vdots \\ \mathbf{\mu}_d \end{bmatrix} = \begin{bmatrix} \mathbf{y}_1 \\ \vdots \\ \mathbf{y}_K \\ \mathbf{\mu}_d \end{bmatrix}$$

vector of observations

$$\begin{bmatrix} \mathbf{y}_1 \\ \vdots \\ \mathbf{y}_K \\ \mathbf{\mu}_d \end{bmatrix} + \begin{bmatrix} \mathbf{e}_1 \\ \vdots \\ \mathbf{e}_K \\ \mathbf{e}_d \end{bmatrix} = \begin{bmatrix} \mathbf{b}_{J+1,1}^T \\ \vdots \\ \mathbf{b}_{J+1,k}^T \\ \mathbf{I} \end{bmatrix} \mathbf{d}_J$$

($N \times n_k$) ($N \times 1$)

matrix of scal. functions vector of scal. coefficients

$$D \left(\begin{bmatrix} \mathbf{y}_1 \\ \vdots \\ \mathbf{\mu}_d \end{bmatrix} \right) = \sigma_k^2 \begin{bmatrix} \mathbf{P}_1^{-1} & \mathbf{0} & \dots & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ \vdots & \vdots & \ddots & \vdots \\ \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \end{bmatrix} + \dots + \sigma_d^2 \begin{bmatrix} \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \dots & \mathbf{0} \\ \vdots & \vdots & \ddots & \vdots \\ \mathbf{0} & \mathbf{0} & \dots & \Sigma_d^{-1} \end{bmatrix}$$

$k = 1 \dots K$ various observation techniques

n_k number of observations from technik k

Note: the measurements \mathbf{y} are treated as independent observations, i. e. without correlations.

- Solving the normal equations (by iteratively determined VCs) results in $\hat{\mathbf{d}}_J$.
- Extracting the erroneous observations \mathbf{y} and applying the law of error propagation then results in the variance covariance matrix $D(\hat{\mathbf{d}}_J)$.

Detail signal $j = 11$

G_{11}	min ... max [mGal]	mean [mGal]	+/- std. [mGal]
Dg_{11}	-23.86 ... 39.74	0.06	3.58
σ_{11}	0.11 ... 18.18	0.86	1.61

large standard deviations in areas of replaced
„mean-coefficients“

-> outlier detection $|Dg_{11}| \leq 3\sigma_{11}$

$Dg_{11, \text{red}}$	-17.44 ... 12.64	-0.07	2.96
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