

Height Estimation and Error Assessment of Inland Water Level Time Series calculated by a Kalman Filter Approach using Multi-Mission Satellite Altimetry

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In this poster, we present our methodology for estimating water level time series over lakes, reservoirs, and rivers. Furthermore, an error assessment of the resulting water level time series is demonstrated. For computing the water level time series multimission satellite altimetry data is used. The estimation is based on altimeter data from Topex, Jason-1, Jason-2, GFO, ERS-2, Envisat, and Saral/Altika – depending on the location of the water body. According to the extent of the investigated water body 1Hz, high-frequent or retracked altimeter measurements can be used.

Different outlier rejections such as thresholds or Support Vector Regression (SVR) can be used for rejecting outliers. For estimating water level heights a Kalman filter approach applied to the grid nodes of a hexagonal grid covering the water body of interest is implemented. After setting an error limit on the resulting water level heights of each grid node, a weighted average water level per point of time is derived referring to one reference location.

For the estimation of water level height accuracies, at first, the formal errors are computed applying a full error propagation within Kalman filtering. Hereby, the precision of the input measurements are introduced by using the standard deviation of the water level height along the altimeter track.

For validation of the time series, we compare our results with in situ data and external inland altimeter databases from Hydroweb (LEGOS), River and Lakes (ESA-DMU), and Global River and Lakes Monitor (GRLM) by USDA. We yield low RMS differences and very high correlations between absolute water level height time series from altimetry and gauges. Moreover, the comparisons of water level heights are also used for the validation of the error assessment.

More than 250 water level time series were already computed and made public available via the "Database for Hydrological Time

Database for Hydrological Time Series of Inland Water (DAHITI)

The resulting water level time series over inland water are accessible via the website of the "Database for Hydrological Time Series of Inland Water (DAHITI)" (Schwatke et al., 2014). The DAHITI web-service is operated by the Deutsches Geodätisches Forschungsinstitut der Technischen Universität München (DGFI-TUM) and is available under http://dahiti.dgfi.tum.de. The data holding of DAHITI contains about 250 time series over lakes, rivers and wetlands. The water level time series can be downloaded for free after a short registration process.

http://dahiti.dgfi.tum.de

Figure 4: Website of the "Database for Hydrological Time Series of Inland Water " (DAHITI) at DGFI-TUM



Data

The data holding for the estimation of water level time series is the **Open Altimeter Database (OpenADB)** (Schwatke et al. 2014a) of the DGFI-TUM. Hereby, satellite altimeter data from all available altimeter missions can be used depending on the data coverage of the investigated inland water body. All altimeter data are cross-calibrated to remove the range bias between the missions to allow the usage of all missions as a single altimeter system.



Methodology

The procedure for the estimation of water level time series is divided into three steps: "Preprocessing", "Kalman filtering" and "Postprocessing" which are described as follows.

Preprocessing

• Extraction of altimeter data from OpenADB for the investigated inland water body Optional retracking of the altimeter measurements using the Improved Threshold retracker (Hwang et al., 2006) with a threshold of 10% Calculation of final normal heights for each measurement by applying geophysical corrections Calculation of standard deviations of the heights along each altimeter track based on a moving window (0.5 km for rivers, 1.5 km for lakes) User defined outlier detection based on thresholds for the min./max. latitude, min./max. height or max. standard deviation can be applied. An additional outlier detection using a linear Support Vector Regression (Smola et al, 2004) can be applied to reject water level heights not representing a flat surface.



Results and Validation

For the validation of water levels from DAHITI comparisons with in-situ gauges from the *Agência Nacional de Águas* and the *Canada Wateroffice* are performed. Furthermore, to show the potential of the Kalman filter approach, water level time series from LEGOS (Cretaux et al., 2011), ESA-DMU (Berry et al, 1997), GRLM (Birkett et al., 2011) are also compared with in-situ data.

The water level time series of Lake Athabasca (7,800 km²) in Canada is presented in detail. Fig. 3 shows the water level time series from DAHITI, LEGOS, ESA-DMU and GRLM compared with in-situ data and shifted to the water level height of DAHITI. Furthermore, The year 2004 is plotted in detail. The comparison between DAHITI and in-situ data shows a very good agreement with RMS differences of 17cm and a correlation coefficient of 0.88. Furthermore, we can increase the temporal resolution and up to 1337 points which are used for the due to the multi-mission approach. The comparison between in-situ data and LEGOS (RMS=33.7cm, R²=0.79, 272 points), ESA-DMU (RMS=80.5cm, R²=0.30, 79 points), and GRLM (RMS=55.7cm, R²=0.27, 76 points) show higher RMS differences and smaller correlation. Fig.3 also clearly show the improvement of the water level time series in parts of ice coverage in winter compared to the other time series from altimetry.



For a further validation comparisons between DAHITI and in-situ data and time series from other groups for selected lakes in Canada and rivers in the Amazon basin are performed. Table 1 shows the RMS differences, correlation

Kalman Filtering

- Estimation of a hexagonal computation grid over the inland water body
- Resulting heights and standard deviations from the preprocessing step are used as input data for the Kalman filtering.
- The method of Kalman Filter (Kalman, 1960) is applied. Hereby, water level heights and formal errors for every epoch in which new altimeter measurements are introduced and for every grid node over the water body are computed. The applied Kalman filter approach contains a rigorous error propagation.

Postprocessing

- For every epoch a final water level height and a formal error are estimated
- Finally, a last outlier detection can be applied on the water level time series based on SVR using radial base functions to





Figure 2: Processing strategy for the computation of water level time series (Schwatke et al., 2015)

Error estimation

The current error estimation of the Kalman filter only yields formal errors which are too optimistic. At the moment, there are no uncertainties of geophysical corrections or altimeter ranges considered. Furthermore, no correlation between adjacent altimeter measurements are considerd. The improvement of the resulting errors are still

coefficients and number of points used for the.

Target name	Station name	Retracked	Ice Coverage	DAHITI			LEGOS			ESA-DMU			GRLM		
				RMS [cm]	R ²	No	RMS [cm]	R²	No	RMS [cm]	R²	No	RMS [cm]	R ²	No
Lake Athabasca	Crackingstone Point ¹	Х	Х	17.0	0.88	1337	33.7	0.79	272	80.5	0.30	79	55.7	0.27	76
Great Slave Lake	Hay River ¹		Х	15.9	0.61	1221	31.2 0.37 246 n.a.		n.a.						
Lake Claire	Prairie Point ¹	Х	Х	19.9	0.53	141	n.a.			37.9	0.25	70	n.a.		
Lake Winnipeg	George Island ¹ Gimli ¹ Pine Dock ¹		X X X	14.6 24.4 19.1	0.81 0.49 0.65	522 518 499	27.7 30.0 28.3	0.64 0.55 0.63	257 264 254	41.9 41.9 42.7	0.49 0.48 0.51	77 76 74	30.9 35.9 33.8	0.56 0.42 0.50	703 722 702
Lake Manitoba	Westbourne ¹ Steep Rock ¹	X X	X X	20.0 20.0	0.77 0.77	318 315	34.2 36.4	0.42 0.40	100 101	34.2 35.5	0.42 0.33	73 75	46.0 47.3	0.11 0.11	71 72
Cedar Lake	Oleson Point ¹		Х	37.8	0.83	530	76.7	0.20	252	54.8	0.54	78	n.a.		
Lake Winnipegosis	Winnipegosis ¹	Х	Х	16.2	0.91	426	36.7	0.63	136	34.2	0.61	70	n.a.		
Lake of the Woods	Clearwater Bay ¹ Cyclone Island ¹ Hanson Bay ¹	X X X	X X X	16.3 15.0 14.7	0.75 0.78 0.79	631 631 631	44.4 43.4 42.6	0.58 0.61 0.63	143 143 143	36.6 35.9 36.0	0.40 0.41 0.40	77 77 77	n.a. n.a. n.a.		
Rio Solimoes	Tefé ² Tefé ² Itepé ² ua Itepéua ²	X X X X		11.6 21.7 49.9 46.0	1.00 0.99 0.98 0.99	32 72 39 117	53.9 120.4 61.2	n.a. 0.98 0.90 0.97	84 44 97	173.3	<mark>0.76</mark> n.a. n.a. n.a.	27	n.a. n.a. n.a. n.a.		
Rio Purus	Aruma-Jusante ²	Х		23.3	1.00	16	24.1	1.00	7	318.9	0.61	6	n.a.		
Rio Madeira	Humanitá ² Humanitá ² Guajará-Mirim ²	X X X		21.6 29.0 78.4	1.00 1.00 0.88	35 93 44	45.1 50.2 87.4	0.99 0.99 0.88	29 <mark>91</mark> 35	53.2 134.3	0.99 n.a. 0.77	28 36	n.a. n.a. n.a.		
Rio Negro	Porto de Manaus ² Moura ² Moura ²	X X X		12.7 65.3 46.0	1.00 0.97 0.98	74 62 42	21.8 70.9 46.3	1.00 0.96 0.97	72 43 45	67.5 44.1	0.96 n.a. 0.98	71 40	n.a. n.a. n.a.		

Table 1: Water level time series of DAHITI for selected lakes and rivers compared with in-situ data and results from LEGOS, ESA-DMU, and GRLM (Source of in-situ data: ¹Canada Wateroffice, ²Agência Nacional de Águas) (Schwatke et al., 2015)

Discussion/Outlook

- DAHITI provides time series over inland water for hydrological applications
- Comparisons of results from DAHITI with in-situ data over selected lakes and river show a decrease of RMS differences and an increase of the correlation coefficients compared with water level time series from LEGOS, ESA-DMU and GRLM
- The new approach uses a Kalman filter and leads to reliable and high accurate water level time series over inland water

under investigation.

DAHITI provides formal errors, which represent the temporal variation of errors very well but at the moment, the resulting errors are too optimistic due to the used standard deviation as error for the Kalman filter approach
In future, the Kalman filter approach can also be used for the prediction of water level time series over inland water

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